

# Study of Heat and Mass Transport in Bénard-Darcy Convection with G-Jitter and Variable Viscosity Liquids in a Porous Layer with Internal Heat Source

A. Srivastava<sup>1†</sup>, B. S. Bhadauria<sup>2</sup>, and A. K. Singh<sup>1</sup>

<sup>1</sup> *Department of Mathematics, Institute of Science, Banaras Hindu University, Varanasi-221005, India*

<sup>2</sup> *Department of Applied Mathematics, School for Physical Sciences, Babasaheb Bhimrao Ambedkar University, Lucknow 226025, India*

† *Corresponding Author Email: srivastavaalok0311@gmail.com*

(Received September 29, 2017; accepted May 4, 2018)

## ABSTRACT

In this research article, we investigated the weakly non-linear effect of gravity modulation for the temperature dependent viscous fluid in a horizontal porous layer in the presence of internal heat source. We use power series expansion in terms of the amplitude of gravity modulation, which is considered to be small for double-diffusive convection in porous media. We graphically show the effect of internal heat source, solute Rayleigh number, Lewis number, Vadász number, thermo-rheological parameter, the amplitude of gravity modulation, the frequency of modulation on the heat and mass transfer using Ginzburg-Landau equation. The effect of gravity modulation is found significant and is more effective for the low values of frequency of modulation.

**Keywords:** Ginzburg-Landau equation; Gravity modulation; Porous media; Temperature dependent viscosity; Double-diffusive convection; Weakly non-linear stability.

## NOMENCLATURE

### Latin Symbols

$A$	amplitude of convection
$a_1$	amplitude of gravity modulation
$d$	depth of the fluid layer
$\mathbf{g}$	acceleration due to gravity
$\mathbf{K}$	permeability of the porous medium
$k_c$	critical wave number
$Le$	Lewis number, $Le = \frac{\kappa_T}{\kappa_S}$
$Nu$	Nusselt number
$p$	reduced pressure
$Ra_T$	thermal Rayleigh number, $Ra_T = \frac{\beta_T g_0 \Delta T d K}{\nu \kappa_T}$
$Ra_S$	solute Rayleigh number, $Ra_S = \frac{\beta_{Sg_0} \Delta S d K}{\nu \kappa_T}$
$R_i$	internal heat source parameter $R_i = \frac{Qd^2}{\kappa_T}$
$S$	solute concentration
$\Delta S$	solute difference across the porous layer
$Sh$	Sherwood number
$T$	temperature
$\Delta T$	temperature difference across the porous layer
$t$	time

$\beta_T$	coefficient of thermal expansion
$\gamma$	heat capacity ratio $\gamma = \frac{(\rho c)m}{(\rho c)f}$
$\delta_0$	small parameter variation of viscosity with temperature
$\varepsilon$	perturbation parameter
$\kappa_S$	solutal diffusivity
$\kappa_T$	thermal diffusivity
$\mu$	dynamic viscosity of the fluid
$\nu$	kinematic viscosity, $\left(\frac{\mu}{\rho_0}\right)$
$\rho$	fluid density
$\tau$	slow time $\tau = \varepsilon^2 t$
$\phi$	porosity
$\psi$	stream function
$\Omega$	frequency of modulation

### Other symbols

$$\nabla_1^2 = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2}$$

$V$	thermorheological parameter $V = \delta_0 \Delta T$
$Va$	Vadász number $Va = \frac{vd^2}{K\kappa_T}$
$(x,z)$	horizontal and vertical co-ordinates

**Greek Symbols**

$\beta_S$	coefficient of solute expansion
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$$\nabla^2 = \nabla_1^2 + \frac{\partial^2}{\partial z^2}$$

**Subscripts**

$b$	basic state
$c$	critical
$0$	reference value

**Superscripts**

'	perturbed quantity
*	dimensionless quantity
$St$	stationary

**1. INTRODUCTION**

Horton-Rogers-Lapwood convection attracts many researchers in recent years as it come across many natural convection and in most geological, physical, mechanical engineering, chemical engineering, geophysical problems. Documented review in this area are mainly provided by [Ingham and Pop \(2005\)](#), [Vadász \(2008\)](#), [Nield and Bejan \(2013\)](#), [Vafai \(2005\)](#). Double-diffusive convection in which two components having different rates of diffusivity assent to the instability arises in oceans, where heat and salt concentration exist with different gradients and diffuse at differing rates, in our atmosphere where temperature and concentrations have different diffusive rates at different altitudes with respect to the sea level, magma chambers, polymeric liquids, geothermal energy extraction, oil re-recovery process, nuclear waste disposal and the migration of moisture through air contained in fibrous insulation. [Kuznetsov and Nield \(2008\)](#), [Kuznetsov and Nield \(2010\)](#), [Kuznetsov and Nield \(2011\)](#), [Bhadoria \(2012\)](#) undertook the thermal instability of double diffusive convection in porous media.

Under the micro-gravity environment gravity field is a randomly fluctuating field and thus the gravity modulation of the system leads to the variable coefficients in the governing equations of thermal instability in porous media and involves the vertical time-periodic vibrations of the system. This leads to the appearance of a modified gravity, collinear with actual gravity, in the form of a time-periodic gravity field perturbation and is known as gravity modulation or g-jitter in literature. Gravity modulation can be taken as an effective mechanism to control the heat transfer by an external regulation. Documented article in this area are given by [Malashetty and Padmavathi \(1997\)](#), [Rees and Pop \(2000\)](#), [Rees and Pop \(2001\)](#), [Rees and Pop \(2003\)](#), [Govender \(2005\)](#), [Kuznetsov \(2005\)](#), [Kuznetsov \(2006b\)](#), [Kuznetsov \(2006a\)](#), [Strong \(2008\)](#), [Strong \(2009\)](#), [Razi \*et al.\* \(2009\)](#), [Vanishree and Siddheshwar \(2010\)](#), [Siddheshwar \*et al.\* \(2012a\)](#), [Swamy \*et al.\* \(2013\)](#), [Swamy \(2014\)](#).

There are a large number of practical situations in which convection is driven by the internal heat source. Due to internal heating of earth, there is a temperature gradient between the interior and the

exterior of the earth's crust which causes convection in earth crust also application of internal heat source may be found in radioactive decay of un-stable isotopes, metal waste form development for spent nuclear fuel, weak exothermic reaction which can take place within porous materials moreover internal heat source is the main energy source of celestial bodies which is generated by radioactive decay and nuclear reaction. Related research article in this area is provided by [Tveitereid \(1977\)](#), [Bejan \(1978\)](#), [Alex \*et al.\* \(2001\)](#), [Saravanan \(2009\)](#), [Cookey \*et al.\* \(2010\)](#), [Nouri-Borujerdi \*et al.\* \(2007\)](#), [Nouri-Borujerdi \*et al.\* \(2008\)](#), [Capone \*et al.\* \(2011\)](#), [Bhadoria \(2012\)](#), [Bhadoria \*et al.\* \(2013\)](#).

The viscosity of a fluid is one of the property which varies as temperature varies but most of the articles kept it as a constant with respect to temperature, for some fluids it is less and for some fluids it is significant. Temperature dependent viscosity fluid gives rise to variation in the top and bottom structures and referred as a non-Boussinesq effect. Nonlinear energy stability theory has been derived by [Richardson and Straughan \(1993\)](#) for the problem of convection in porous medium when the viscosity depends on the temperature for vanishingly small initial data thresholds, [Payne \*et al.\* \(1999\)](#) has been studied the unconditional nonlinear stability for temperature sensitive fluid in porous media, further more [Payne and Straughan \(2000\)](#) extend their analysis for the cases when the viscosity variation may be quadratic or when convection is penetrative. [Qin and Chadam \(1996\)](#) studied the nonlinear energy stability by considering the temperature dependent viscosity and inertial drag by taking higher-order approximations for the viscosity-temperature and density-temperature relation, [Nield \(1996\)](#), [Holzbecher \(1998\)](#) studied the thermal instability for the variable viscosity fluid in porous layer using the FAST-C(SD) code for numerical modeling, [Rees \*et al.\* \(2002\)](#), [Siddheshwar and Chan \(2004\)](#), [Vanishree and Siddheshwar \(2010\)](#) investigated the linear thermal instability for the temperature dependent fluid using weighted residual technique, [Siddheshwar \*et al.\* \(2012b\)](#) studied the effects of variable viscosity and the gravity modulation on the heat transfer in an anisotropic porous medium, [Srivastava \*et al.\* \(2013\)](#) analyse the non-linear effect of internal heat source and gravity modulation on the heat

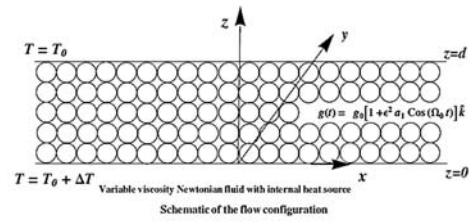
transfer for variable viscosity fluid in an anisotropic porous layer.

The simultaneous appearance of solid and liquid phases during the solidification of metallic alloys referred as mushy zones (Worster (1991)), also during the solidification of metals there is a temperature difference between upper and lower layer of molten metals which solidify slowly and in which the heat consumed by the molten metals works as the internal heat source. In the present paper the study of extraction of metals from ores where a mushy layer is formed during solidification of a metallic alloy is of particular interest. The quality and structure of the resulting solid during the solidification of binary alloys can be controlled by influencing the transport process externally, which can be done by thermal modulation, gravity modulation, and rotation or by internal heating. However, in the present study, internal heating and gravity modulation of the system was used as an external means to influence the transport process, thereby controlling the quality and structure of the resulting solid. If one were to quantify heat and mass transports in porous media in the presence of gravity modulations, then the linear stability analysis is inapplicable and the nonlinear stability analysis becomes inevitable. Numerical analysis of the Ginzburg-Landau equation is simpler than numerical solution of the equations of motion. In addition, analysis of stability of some simple (for example, periodic) solutions of the Ginzburg-Landau equation allows researchers to simplify the analysis of spatiotemporal dynamics of complex flows in fluid mechanics, while investigating a weakly non-linear stability of systems Ginzburg-Landau equation arise as a consequence of the solvability condition in a large category of problems in continuum mechanics. In the light of the above, we make a weakly nonlinear analysis of the problem using the Ginzburg-Landau equation and, in the process, quantify the heat and mass transports in terms of the amplitude governed by the Ginzburg-Landau equation. There are no reported studies on this aspect of the problem.

## 2. GOVERNING EQUATIONS

We consider an infinitely extended horizontal porous layer saturated by variable viscosity Newtonian fluid with temperature-dependent viscosity confined between the planes  $z = 0$  and  $z = d$ , which is heated and salted from below. We choose Cartesian frame of reference as, origin in the lower boundary and the  $z$ -axis vertically upward direction. The gravity force is acting in vertically down-ward direction, we consider only free-free boundaries. A uniform adverse temperature gradient  $\Delta T$  and concentration gradient  $\Delta S$  is maintained between the surfaces. Further the density variation is considered under Boussinesq approximation. The governing equations under above considerations are given by

$$\nabla \cdot \mathbf{q} = 0, \tag{1}$$



$$\frac{\rho_0}{\phi} \frac{\partial \mathbf{q}}{\partial t} = -\nabla p + \rho \mathbf{g}(t) - \mu \mathbf{K} \cdot \mathbf{q}, \tag{2}$$

$$\gamma \frac{\partial T}{\partial t} + (\mathbf{q} \cdot \nabla) T = \kappa_T \nabla^2 T + Q(T - T_0), \tag{3}$$

$$\phi \frac{\partial S}{\partial t} + (\mathbf{q} \cdot \nabla) S = \kappa_S \nabla^2 S, \tag{4}$$

$$\rho = \rho_0 [1 - \beta_T (T - T_0) + \beta_S (S - S_0)], \tag{5}$$

$$\mathbf{g}(t) = g_0 [1 + \epsilon^2 a_1 \cos(\Omega_0 t)] \hat{k}, \tag{6}$$

$$\mu(T) = \frac{\mu_0}{1 + \epsilon^2 \delta_0 (T - T_0)}, \tag{7}$$

where,  $\mathbf{q} = (u, v, w)$  is velocity (m/s),  $\phi$  is porosity of the matrix,  $p$  is the pressure (Pa),  $\mathbf{g}$  is the acceleration due to gravity ( $\text{m/s}^2$ ),  $\mu$  is viscosity ( $\text{Ns/m}^2$ ) and is taken as in the sense that  $1/\mu$  is linear in the temperature because the steady state temperature profile in the Horton-Rogers-Lapwood problem in the vertical coordinate  $z$  is linear, this gives a situation where  $1/\mu$  is linear in  $z$ , the ordinary differential equation that now arises is one whose coefficients are linear in  $z$  and this allow us to solve the equation using a comparatively simple series method.  $\rho$  is density ( $\text{kg/m}^3$ ),  $T$  and  $S$  represents temperature (K) and concentration ( $\text{kg/m}^3$ ) respectively.  $\mathbf{K}$  is the inverse of the permeability tensor ( $\text{m}^2$ ),  $\kappa_T$  is the thermal diffusivity tensor ( $\text{m}^2/\text{s}$ ),  $Q = Q_0 / (\rho c) f$  and  $Q_0$  is the volumetric internal heat source ( $\text{W/m}^3$ ) and the magnitude of the internal heat generation can vary with the temperature of the elements,  $\rho_0$  is reference density,  $g_0$  is mean gravity,  $a_1$  is amplitude of gravity modulation,  $\Omega_0$  is the frequency ( $\text{s}^{-1}$ ),  $\epsilon$  is the quantity that indicates smallness in order of magnitude of modulation and  $t$  is time. Furthermore  $\kappa_S$  is the solute diffusivity ( $\text{m}^2/\text{s}$ ),  $\beta_T$  is thermal expansion coefficient ( $\text{K}^{-1}$ ) and  $\beta_S$  is density coefficient for salinity ( $\text{kg}^{-1} \text{m}^3$ ),  $\gamma = \frac{(\rho c) m}{(\rho c) f}$  is the heat

capacity ratio where  $c$  is the specific heat (J/kg K) and subscript  $f$  and  $m$  stands for fluid and medium, also for simplicity  $\gamma$  and  $\phi$  is taken unity in this paper. We consider only two-dimensional disturbances in our study, introducing the stream function  $\psi$  and eliminating the pressure term and then nondimensionalizing the resultant equations using the substitution

$$(x, y, z) = (x^*, y^*, z^*)d, t = t^*(\gamma d^2 / \kappa_T),$$

$$\psi = (\kappa_T)\psi^*, T - T_0 = (\Delta T)T^*, S = (\Delta S)S^* \quad (8)$$

The nondimensionlized (after dropping the asterisks for simplicity) Eqs. (2-4) are:

$$\frac{1}{Va} \frac{\partial(\nabla^2 \psi)}{\partial t} = -Ra_T(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0 t)) \frac{\partial T}{\partial x}$$

$$+ Ra_S(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0 t)) \frac{\partial S}{\partial x}$$

$$- \bar{\mu}(T) \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} \right) \psi$$

$$- \left( \frac{\partial \bar{\mu}}{\partial x} \frac{\partial \psi}{\partial x} + \frac{\partial \bar{\mu}}{\partial z} \frac{\partial \psi}{\partial z} \right), \quad (9)$$

$$\frac{\partial T}{\partial t} - \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} + R_i \right) T = \frac{\partial(\psi, T)}{\partial(x, z)}, \quad (10)$$

$$\frac{\partial S}{\partial t} - \frac{1}{Le} \nabla^2 S = \frac{\partial(\psi, S)}{\partial(x, z)}, \quad (11)$$

where  $\bar{\mu}(T) = \frac{1}{1 + \varepsilon^2 V(T - T_0)}$  and the

appearance of  $\varepsilon^2$  indicates that the viscosity variation is weak as  $\varepsilon$  is small quantity.

$Va = \frac{\nu d^2}{K \kappa_T}$  is the Vadász number or Darcy-

Prandtl number  $Ra_T = \frac{\beta_T g \Delta T d K}{\nu \kappa_T}$  is thermal

Rayleigh number,  $Ra_S = \frac{\beta_S g \Delta S d K}{\nu \kappa_T}$  is solute

Rayleigh number,  $R_i = \frac{Q d^2}{\kappa_T}$  Internal heat source

parameter,  $Le = \frac{\kappa_T}{\kappa_S}$  is Lewis number,  $V = \delta_0 \Delta T$

thermo-rheological parameter.

The boundary conditions for solving Eqs. (9-11) are

$$\psi = 0, T = 1 \text{ and } S = 1 \text{ at } z = 0 \quad (12)$$

$$\psi = 0, T = 0 \text{ and } S = 0 \text{ at } z = 1. \quad (13)$$

The boundaries of a porous medium can be either free or rigid. We know that under laboratory conditions free boundaries are less accessible to the experiments therefore; one has to consider the rigid boundaries. However, in real situations like geothermal regions the porous layer under consideration cannot be isolated from the surrounding region to avoid the penetration of the fluid; therefore we have to consider only free surfaces. Further, the mathematical analysis of the problem becomes easier to handle, therefore for mathematical convenience also the free-free boundary conditions are used.

The conduction profile is given by

$$\psi_b = 0, T_b(z) = \frac{\text{Sin} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i}} \quad (14)$$

and  $S_b(z) = 1 - z$

Using  $\bar{\mu} = \bar{\mu}(T_b)$  Eq. (9) reduces to

$$\frac{1}{Va} \frac{\partial(\nabla^2 \psi)}{\partial t} = -Ra_T(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0 t)) \frac{\partial T}{\partial x}$$

$$+ Ra_S(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0 t)) \frac{\partial S}{\partial x}$$

$$- \bar{\mu}(T) \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} \right) \psi$$

$$- \frac{\partial \bar{\mu}}{\partial z} \frac{\partial \psi}{\partial z},$$

where  $\bar{\mu}(T_b) = \frac{\mu'}{\mu_0} = \frac{1}{\left[ 1 + \varepsilon^2 V \frac{\text{Sin} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i}} \right]}$

and  $\Omega_0^* = \frac{\Omega_0 d^2}{\kappa_{Tz}}$ .

Now imposing finite amplitude perturbations on the basic quiescent state given by Eq. (14) as

$$\psi = \Psi, T = \frac{\text{Sin} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i}} + \Theta \quad (16)$$

and  $S = 1 - z + \phi$

Substituting the above Eq. (16) in Eqs. (10, 11, 15) we have

$$\frac{1}{Va} \frac{\partial(\nabla^2 \psi)}{\partial t} = -Ra_T(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0^* t)) \frac{\partial \Theta}{\partial x}$$

$$+ Ra_S(1 + \varepsilon^2 a_1 \text{Cos}(\Omega_0^* t)) \frac{\partial \phi}{\partial x}$$

$$- \left[ \frac{1}{1 + \varepsilon^2 V \frac{\text{Sin} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i}}} \right] \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} \right) \psi \quad (17)$$

$$- \frac{\varepsilon^2 V \sqrt{R_i} \text{Cos} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i} \left[ 1 + \varepsilon^2 V \frac{\text{Sin} \sqrt{R_i} (1 - z)}{\text{Sin} \sqrt{R_i}} \right]^2} \frac{\partial \psi}{\partial z},$$

$$- \frac{dT_b}{dz} \frac{\partial \psi}{\partial x} + \frac{\partial \Theta}{\partial t}$$

$$- \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} + R_i \right) \Theta = \frac{\partial(\psi, \Theta)}{\partial(x, z)}, \quad (18)$$

$$\frac{\partial \psi}{\partial x} + \frac{\partial \phi}{\partial t} - \frac{1}{Le} \nabla^2 \phi = \frac{\partial(\psi, \Theta)}{\partial(x, z)} \quad (19)$$

Boundary conditions to solve Eqs. (17-19) are

$$\psi = 0, \Theta = 0 \text{ and } \phi = 0 \text{ at } z = 0, \quad (20)$$

$$\psi = 0, \Theta = 0 \text{ and } \phi = 0 \text{ at } z = 1. \quad (21)$$

We now introduce the following asymptotic expansion

$$Ra_T = R_{0c} + \varepsilon^2 R_2 + \varepsilon^4 R_4 + \dots, \quad (22)$$

$$\psi = \varepsilon \psi_1 + \varepsilon^2 \psi_2 + \varepsilon^3 \psi_3 + \dots, \quad (23)$$

$$\Theta = \varepsilon \Theta_1 + \varepsilon^2 \Theta_2 + \varepsilon^3 \Theta_3 + \dots, \quad (24)$$

$$\phi = \varepsilon \phi_1 + \varepsilon^2 \phi_2 + \varepsilon^3 \phi_3 + \dots, \quad (25)$$

where  $R_{0c}$  is the critical value of the Rayleigh number at which the onset of convection takes place in the absence of gravity modulation.

Since the exponential growth of linearised disturbances is proportional to the difference of the Rayleigh number from the critical value therefore, we assume the variation of time only at the slow time scale  $\tau = \varepsilon^2 t$ , and arranging the systems at different order of  $\varepsilon$ .

At the lowest order, we have

$$\begin{pmatrix} \nabla_1^2 & R_{0c} \frac{\partial}{\partial x} & -Ra_S \frac{\partial}{\partial x} \\ -\frac{dT_b}{dz} \frac{\partial}{\partial x} & -(\nabla_1^2 + R_i) & 0 \\ \frac{\partial}{\partial x} & 0 & -\frac{1}{Le} \nabla^2 \end{pmatrix} \begin{pmatrix} \psi_1 \\ \Theta_1 \\ \phi_1 \end{pmatrix} = 0, \quad (26)$$

Solution at the lowest order subject to the boundary conditions (20, 21) are given by

$$\psi_1 = A[\tau] \sin(k_c x) \sin(\pi z), \quad (27)$$

$$\Theta_1 = -\frac{4\pi^2 k_c}{(\delta^2 - R_i)(R_i - 4\pi^2)} A[\tau] \cos(k_c x) \sin(\pi z), \quad (28)$$

$$\phi_1 = \frac{-k_c Le}{\delta^2} A[\tau] \cos(k_c x) \sin(\pi z), \quad (29)$$

where  $\delta^2 = k_c^2 + \pi^2$  is the critical wave number.

The critical value of the Rayleigh number and the corresponding wave number for the onset of stationary convection is calculated numerically and the expression for Rayleigh number is given by

$$R_{0c} = \frac{(\delta^2 - R_i)(4\pi^2 - R_i)(\delta^4 + k_c^2 Ra_S Le)}{4\pi^2 k_c^2 \delta^2} \quad (30)$$

## 2.1 Amplitude Equation And Heat And Mass Transport For Stationary Instability

At the second order, we have

$$\begin{pmatrix} \nabla_1^2 & R_{0c} \frac{\partial}{\partial x} & -Ra_S \frac{\partial}{\partial x} \\ -\frac{dT_b}{dz} \frac{\partial}{\partial x} & -(\nabla_1^2 + R_i) & 0 \\ \frac{\partial}{\partial x} & 0 & -\frac{1}{Le} \nabla^2 \end{pmatrix} \begin{pmatrix} \psi_2 \\ \Theta_2 \\ \phi_2 \end{pmatrix} = \begin{pmatrix} R_{21} \\ R_{22} \\ R_{23} \end{pmatrix}, \quad (31)$$

where

$$R_{21} = 0, \quad (32)$$

$$R_{22} = \frac{2\pi^3 k_c^2}{(\delta^2 - R_i)(R_i - 4\pi^2)} A[\tau]^2 \sin(2\pi z), \quad (33)$$

$$R_{23} = \frac{-k_c^2 Le \pi}{2\delta^2} A[\tau]^2 \sin(2\pi z) \quad (34)$$

The second order solution subject to the boundary condition (20, 21), is given by

$$\psi_2 = 0, \quad (35)$$

$$\Theta_2 = -\frac{2\pi^3 k_c^2}{(\delta^2 - R_i)(4\pi^2 - R_i)^2} A[\tau]^2 \sin(2\pi z), \quad (36)$$

$$\phi_2 = -\frac{k_c^2 Le^2}{8\pi \delta^2} A[\tau]^2 \sin(2\pi z), \quad (37)$$

The horizontally averaged Nusselt number and Sherwood number,  $Nu$  and  $Sh$ , for stationary mode of convection (the mode considered in this problem) is given by :

$$Nu(\tau) = 1 + \frac{\left[ \frac{k_c}{2\pi} \int_0^{k_c} \left( \frac{\partial \Theta_2}{\partial z} \right) dx \right]_{z=0}}{\left[ \frac{k_c}{2\pi} \int_0^{k_c} \left( \frac{dT_b}{dz} \right) dx \right]_{z=0}}, \quad (38)$$

$$Sh(\tau) = 1 + \frac{\left[ \frac{k_c}{2\pi} \int_0^{k_c} \left( \frac{\partial \phi_2}{\partial z} \right) dx \right]_{z=0}}{\left[ \frac{k_c}{2\pi} \int_0^{k_c} \left( \frac{dS_b}{dz} \right) dx \right]_{z=0}}, \quad (39)$$

One can notice here that the gravity modulation is effective at  $O(\varepsilon^2)$  and effects  $Nu(\tau)$  and  $Sh(\tau)$  through  $A[\tau]$  as shown next. Substituting expressions of  $\Theta_2$  and  $\phi_2$  in the above expression (38, 39) and simplifying, we get

$$Nu(\tau) = 1 + \frac{(4\pi^4 k_c^2 \text{Sin} \sqrt{R_i})}{\left(\sqrt{R_i} \text{Cos} \sqrt{R_i} (\delta^2 - R_i) (R_i - 4\pi^2)\right)^2} (A[\tau])^2, \quad (40)$$

$$Sh(\tau) = 1 + \frac{Le^2 k_c^2}{4\delta^2} (A[\tau])^2 \quad (41)$$

At the third order, we have

$$\begin{pmatrix} \nabla_1^2 & R_{0c} \frac{\partial}{\partial x} & -Ra_S \frac{\partial}{\partial x} \\ -\frac{d\Gamma_b}{dz} \frac{\partial}{\partial x} & -(\nabla_1^2 + R_i) & 0 \\ \frac{\partial}{\partial x} & 0 & -\frac{1}{Le} \nabla^2 \end{pmatrix} \begin{pmatrix} \psi_3 \\ \Theta_3 \\ \phi_3 \end{pmatrix} = \begin{pmatrix} R_{31} \\ R_{32} \\ R_{33} \end{pmatrix}, \quad (42)$$

where

$$\begin{aligned} R_{31} = & -\frac{1}{Va} \frac{\partial}{\partial \tau} (\nabla^2 \psi_1) \\ & -V \frac{\text{Sin} \sqrt{R_i} (1-z)}{\text{Sin} \sqrt{R_i}} \left( \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial z^2} \right) \psi_1 \\ & -V \frac{\text{Cos} \sqrt{R_i} (1-z)}{\text{Sin} \sqrt{R_i}} \frac{\partial \psi_1}{\partial z} \\ & - \left( R_{0c} \left( 2V \frac{\text{Sin} \sqrt{R_i} (1-z)}{\text{Sin} \sqrt{R_i}} + a_1 \text{Cos}(\Omega\tau) \right) \right) \end{aligned} \quad (43)$$

$$\begin{aligned} & + R_2 \frac{\partial \Theta_1}{\partial x} \\ & + Ra_S \left( 2V \frac{\text{Sin} \sqrt{R_i} (1-z)}{\text{Sin} \sqrt{R_i}} + a_1 \text{Cos}(\Omega\tau) \right) \frac{\partial \phi_1}{\partial x}, \end{aligned}$$

$$R_{32} = \frac{\partial \psi_1}{\partial x} \frac{\partial \Theta_2}{\partial z} - \frac{\partial \Theta_1}{\partial \tau}, \quad (44)$$

$$R_{33} = \frac{\partial \psi_1}{\partial x} \frac{\partial \phi_2}{\partial z} - \frac{\partial \phi_1}{\partial \tau} \quad (45)$$

$$\text{and } \Omega = \frac{\Omega_0^*}{\varepsilon^2} = \frac{\Omega_0 d^2}{\varepsilon^2 \kappa_{Tz}}$$

Substituting the value of  $\psi_1, \Theta_1, \Theta_2, \phi_1$  and  $\phi_2$  in the above equations to get the expressions of  $R_{31}, R_{32}, R_{33}$ .

Applying the solvability condition for the existence of third order solution,

$$\int_0^1 \int_0^{k_c} \left( R_{31} \hat{\psi}_1 + R_{32} \hat{\Theta}_1 + R_{33} \hat{\phi}_1 \right) dx dz = 0 \quad (46)$$

where  $(\wedge)$  represents adjoint solution of the first-order system, we get the non-autonomous Ginzburg-Landau equation with time periodic coefficients in the form

$$A_1 A'[\tau] + A_2 A[\tau] + A_3 (A[\tau])^3 = 0, \quad (47)$$

where

$$A_1 = \left( \frac{\delta^2}{Va} + \frac{(\delta^4 + k_c^2 Le Ra_S)}{\delta^2 (\delta^2 - R_i)} - \frac{k_c^2 Le^2 Ra_S}{\delta^4} \right), \quad (48)$$

$$\begin{aligned} A_2 = & -\left( 1 + a_1 \text{Cos}[\Omega\tau] \right) \frac{(\delta^4 + k_c^2 Le Ra_S)}{\delta^2} \\ & + a_1 \text{Cos}[\Omega\tau] \frac{k_c^2 Le Ra_S}{\delta^2} \\ & - \frac{\delta^2 V}{\text{Sin}[\sqrt{R_i}]} \left( \frac{4\pi^2 (-1 + \text{Cos}[\sqrt{R_i}])}{\sqrt{R_i} (R_i - 4\pi^2)} \right) \\ & - \frac{2\pi^2 V \sqrt{R_i} (-1 + \text{Cos}[\sqrt{R_i}])}{(4\pi^2 - R_i) \text{Sin}[\sqrt{R_i}]}, \end{aligned} \quad (49)$$

$$A_3 = \left( \frac{\pi^2 k_c^2 (\delta^4 + k_c^2 Le Ra_S)}{\delta^2 (\delta^2 - R_i) (4\pi^2 - R_i)} - \frac{k_c^4 Le^3 Ra_S}{8\delta^4} \right). \quad (50)$$

The Ginzburg-Landau equation given by (47) is a Bernoulli equation and obtaining its analytical solution is difficult due to its non-autonomous nature. Therefore, it has been solved numerically by the in-built function `NDSolve` of Mathematica 7.0, subject to the initial condition  $A[0] = a_0$ , where  $a_0$  is the chosen initial amplitude of convection. In our calculations, we may assume  $R_2 = R_0$  to reduce the parameters by one.

### 3. RESULTS AND DISCUSSIONS

We perform weakly nonlinear analysis of gravity modulation for temperature dependent viscosity fluid in closely packed porous media, therefore, the Darcy model is considered for the governing equation for linear momentum. The work of [Nield \(1996\)](#) has been used for the thermo-rheological relationship of temperature dependent viscosity of the fluid. We investigated the effect gravity modulation and thermo-rheological parameter on heat and mass transport. We consider the effect of gravity modulation to be of order  $O(\varepsilon^2)$  this will provide us only small amplitude vibrations. Such an assumption will help us in obtaining the amplitude equation of convection in a rather simple and elegant manner and is much easier to obtain than in the case of the Lorenz model. In order to study the heat and mass transfer, we need to perform the nonlinear analysis. We use gravity modulation as an external regulation of convection to control the heat and mass transfer as desired by the need. We consider the gravity modulation, internal heat source and temperature dependent viscous fluid for either enhancing or inhibiting convective heat transport as is required by a real application.

Vadász (1998) pointed out that there are some modern porous medium applications, such as mushy layer in solidification of binary alloys and fractured porous medium, where the value of  $Va$  may be considered to be of unity order, therefore the time-derivative term in the present study has been retained. Further this is the reason that we have kept the values of  $Va$  around one in our calculations, and retained the local acceleration term  $\frac{1}{Va} \frac{\partial q}{\partial t}$ .

The values of  $Nu(\tau)$  and  $Sh(\tau)$  are obtained numerically from the expressions of  $Nu(\tau)$  and  $Sh(\tau)$  by using the numerical value of amplitude obtained from the Ginzburg-Landau equation. We use the values to plot the curve for  $Nu_C(\tau)$  and  $Sh_C(\tau)$  versus  $\tau$  and is presented in the Figs. 1–14. Figures 1-7 correspond to heat transfer and Figs. 8-14 correspond to mass transfer. A close observation of Eqs. (40, 41) in conjunction with Eq. (47) reveals that  $Nu(\tau)$  and  $Sh(\tau)$  depends on Lewis number  $Le$ , solute Rayleigh number  $Ras$ , Vadász number  $Va$ , thermo-rheological parameter and amplitude of g-jitter.

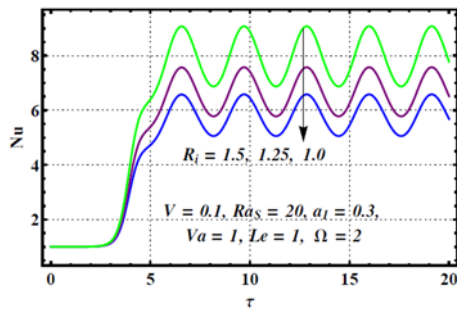


Fig. 1. Variation of Nusselt number with time for different values of  $R_i$

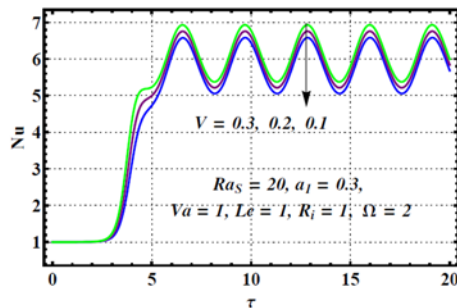


Fig. 2. Variation of Nusselt number with time for different values  $V$

From the Figs. 1–14 it is observed that initially the value  $Nu_C(\tau)$  and  $Sh_C(\tau)$  is one showing that, initially there is no convection and heat transport is prevailed by conduction alone and as time increases heat and mass transport increases shows that convective regime takes place and remains oscillatory for further elapses of time. We take  $\varepsilon = 0.3$  for numerical computations.

From Fig. 1 we observe that to increase in the value of internal heat source parameter the heat

transfer increases it means that if we increase the strength of the internal heat source heat transfer increases. Figure 2 shows the variation of Nusselt number with time for the different values of the thermo-rheological parameter and is clear that to increase in the value

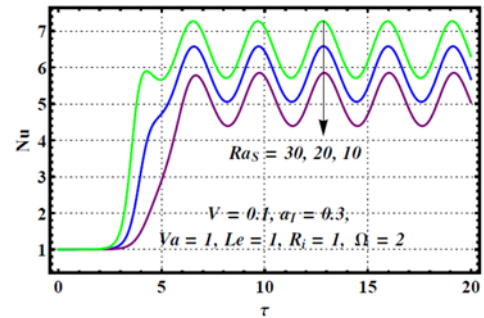


Fig. 3. Variation of Nusselt number with time for different values  $Ras$

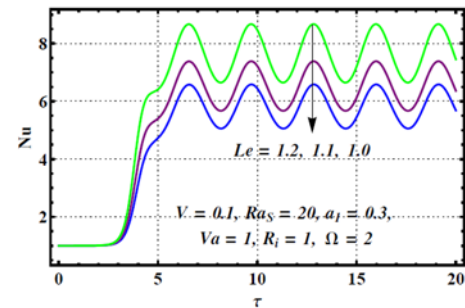


Fig. 4. Variation of Nusselt number with time for different values  $Le$

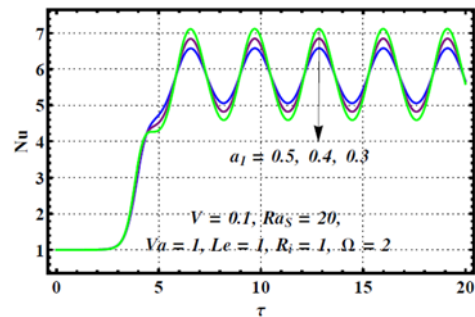


Fig. 5. Variation of Nusselt number with time for different values  $a_1$

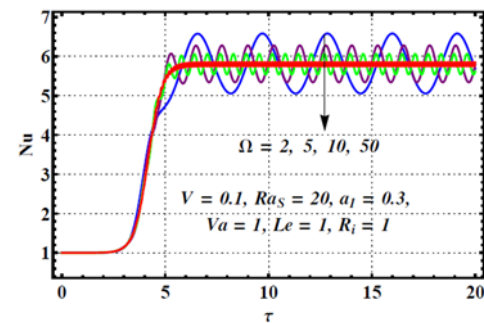


Fig. 6. Variation of Nusselt number with time for different values  $\Omega$

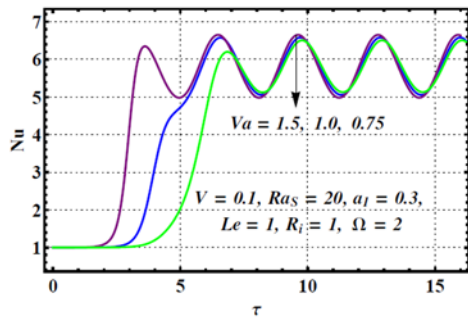


Fig. 7. Variation of Nusselt number with time for different values  $Va$

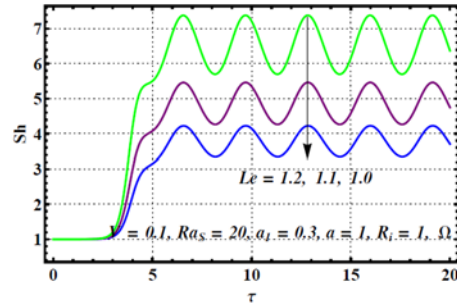


Fig. 11. Variation of Sherwood number with time for different values  $Le$

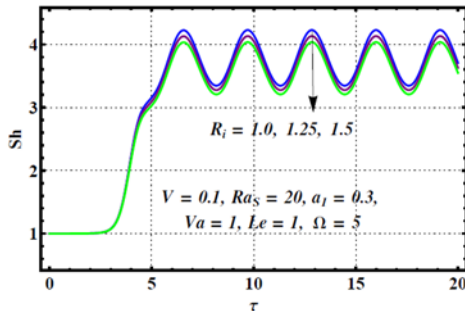


Fig. 8. Variation of Sherwood number with time for different values of  $R_i$

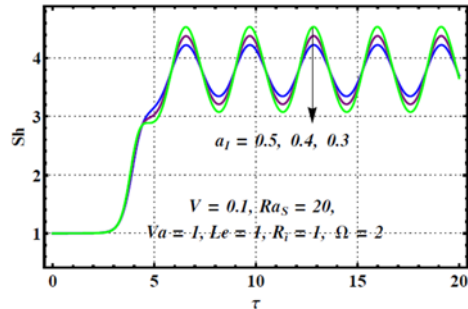


Fig. 12. Variation of Sherwood number with time for different values  $a_1$

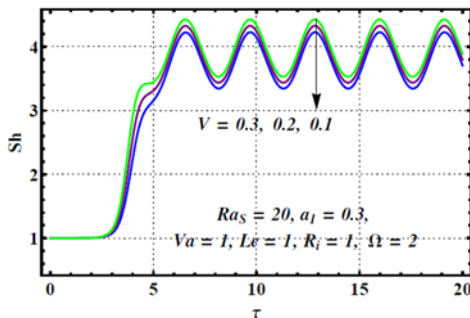


Fig. 9. Variation of Sherwood number with time for different values  $V$

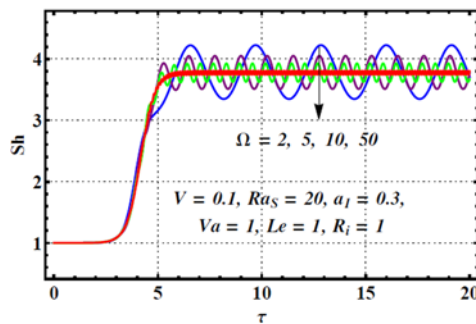


Fig. 13. Variation of Sherwood number with time for different values  $\Omega$

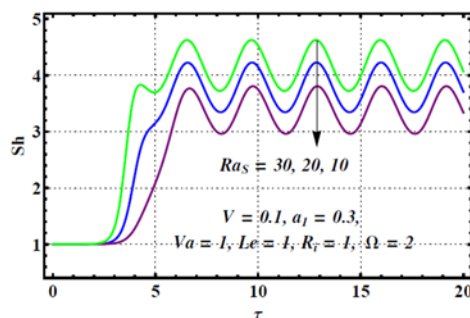


Fig. 10. Variation of Sherwood number with time for different values  $Ra_S$

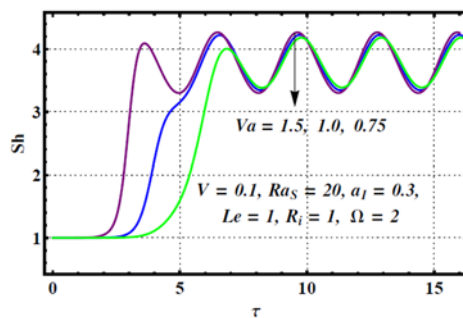


Fig. 14. Variation of Sherwood number with time for different values  $Va$

of thermo-rheological parameter the heat transfer increases this indicates that fluids having high viscosity variation enhancement in the heat transfer is high. Figure 3 reveals that for the increasing value of solute Rayleigh number heat transfer increases that is, if we increase the so-lute gradient the heat transfer increases. Figure 4

shows the effect of Lewis number on the Nusselt number with respect to time and is observed that from the graph that to increase in the Lewis number the heat transfer increases it means that, if we fix the solute diffusivity and increase the thermal diffusivity or decrease the solute diffusivity and fix the thermal diffusivity heat transfer increases. From Fig 5 we observe that to



increase in the amplitude of modulation heat transfer increases it means that if the difference between actual gravity and mean gravity increases heat transfer increases. Figure 6 shows the effect of the frequency of modulation on the heat transfer and it is observed that for the increasing value of frequency of modulation heat transfer decreases this indicates that, if we decrease the time period of the oscillation heat transfer de-creases therefore we can control the heat transfer either to enhance or to inhibit by choosing appropriate frequency of modulation as required by the real life applications, furthermore, for the higher value of the frequency of modulation the effect of gravity modulation become weak this means that for the system having small time period for gravity modulation does not produce significant variation on the heat transfer also the effect of gravity modulation disappears for the higher values of frequency of modulation. From Fig 7 shows that for the increasing value of Vadász number heat transfer in-creases that is, if we increase the Prandtl number and fix the Darcy number or fix the Prandtl number and decrease the Darcy number the heat transfer increases, furthermore, it is observed that Vadász number is more effective on the heat transfer for its lower values so we keep it around unity in our numerical computations.

Figure 8 shows the variation of Sherwood number for the different value of internal heat source parameter with respect to time and is clear from the graph that for the increasing value of internal heat source parameter the mass transfer decreases that is if increase the strength of the heat source mass transfer decreases. From Fig. 9 we observe that to increase in thermo-rheological parameter mass transfer increases it means that for the fluids having viscosity variation high the rate of mass transfer high. From Fig. 10 it is clear that for the increasing value of solute Rayleigh number mass transfer increases this means that if we increase the solute gradient the mass transfer increases. Figure 11 reveals the effect of the Lewis number on the Sher-wood number with respect to time and it is cleat that from the graph that to increase in Lewis number mass transfer increases this indicates that if we fix the solutal diffusivity and increase the thermal diffusivity or decrease the solutal diffusivity and fix the thermal diffusivity mass transfer increases. Figure 12 shows that for the increasing value of the amplitude of modulation mass transfer increases it means that if the difference between actual gravity and mean gravity increases mass transfer in-creases. From Fig. 13 we observe that the effect of frequency of modulation is to decrease the mass transfer for its increasing value this means that if we decrease the time period of the oscillation mass transfer decreases therefore, we can choose the frequency of gravity modulation to control the mass transfer either to enhance or to inhibit as required by the real life applications, furthermore, for the higher value of frequency

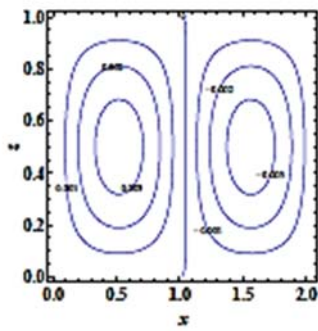
of modulation the effect of gravity modulation is frail on the mass transfer. From Fig. 14 it is clear that to increase in the value of Vadász number mass transfer increases it means that if we increase the Prandtl number and fix the Darcy number or fix the Prandtl number and decrease the Darcy number the mass transfer increases.

Variation of streamlines, isotherms, isohalines at different instant of time is shown graphically in Figs. 15–17. From Figs. 15a–15d it is clear that the magnitudes of streamlines initially increases as time increases and after reaching a certain value it starts oscillations, Fig. 16a–16d shows the variation of isotherms at the different instant of time and if found that from the graph initially isotherms are flat and parallel shows the heat transport is only by conduction and as time increases isotherms starts oscillating showing convective regime is in place and then forms contour showing that as time increases convection contributes in heat transport, similar behaviour is observed for isohalines in Fig. 17a–17d, moreover, it is clear that from the Figs. 15–17 after reaching some instant there is no changes in streamlines, isotherms, isohalines.

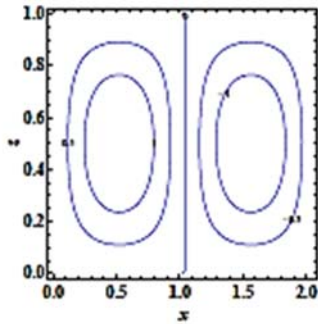
#### 4. CONCLUSION

We perform the weakly nonlinear analysis using the Ginzburg-Landau equation for double diffusive convection in an infinite horizontal porous layer which is heated and salted from below, saturated with temperature sensitive fluid in the presence of gravity modulation and internal heat source. We studied the effect of gravity modulation and temperature dependent viscosity on the heat and mass transfer. We found that the gravity modulation is an effective mechanism to control the heat and mass transfer. The effect of frequency of modulation is significant to control the heat and mass transfer for the physical problem, furthermore for the large value of frequency of modulation the effect of gravity modulation is very frail. We found that the Vadász number is more pronounced on the heat and mass transfer when we take its value around unity, for its higher values the effect of Vadász number is not significant on the heat and mass transfer, however, the effect of gravity modulation keeps its own nature. The following conclusions has been made from our analysis, for the increasing values of parameter:

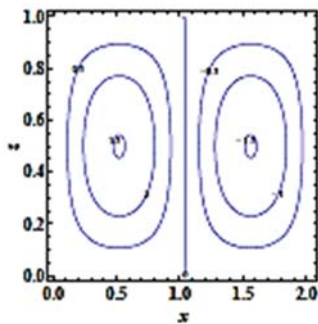
1. Internal heat source parameter  $R_i$  : heat transfer increases, mass transfer decreases.
2. Thermo-rheological parameter  $V$  : heat transfer increases, mass transfer increases.
3. Lewis number  $Le$  : heat transfer increases, mass transfer increases.
4. Solute Rayleigh number  $Ra_s$  : heat transfer increases, mass transfer increases.
5. Vadász number  $Va$  : heat transfer increases, mass transfer increases.



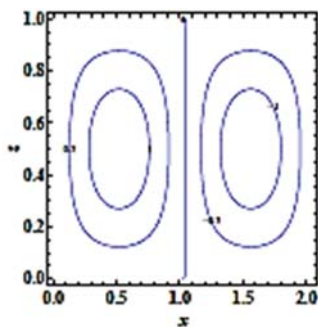
(a)  $\tau = 0.1$



(b)  $\tau = 6$

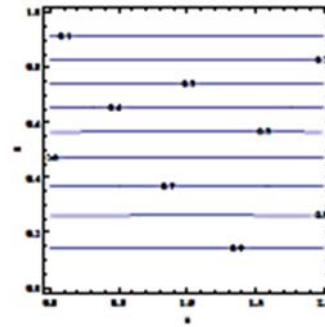


(c)  $\tau = 7$

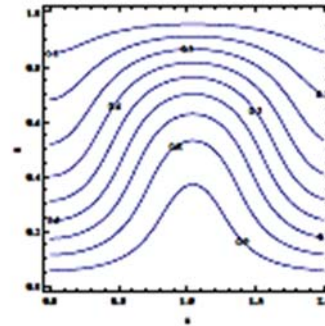


(d)  $\tau = 8$

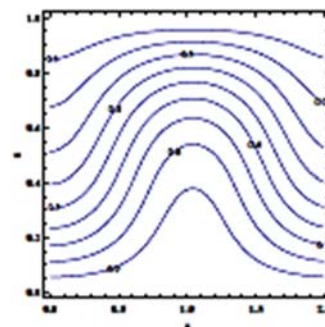
**Fig. 15. Variation of stream lines with time**



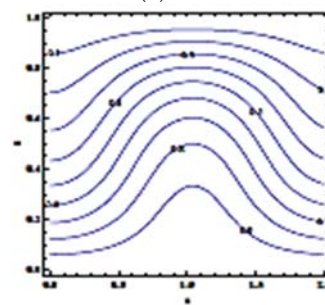
(a)  $\tau = 0.01$



(b)  $\tau = 6$



(c)  $\tau = 7$



(d)  $\tau = 8$

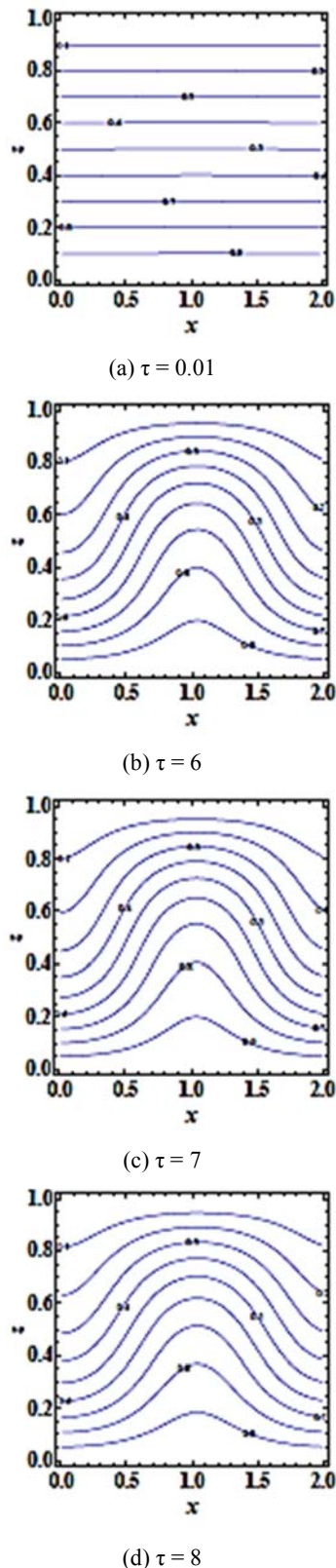
**Fig. 16. Variation of isotherms with time**

6. Frequency of modulation  $\Omega$ : heat transfer decreases, mass transfer decreases.

7. amplitude of modulation  $a_1$ : heat transfer increases, mass transfer increases.

#### ACKNOWLEDGMENTS

Author Alok Srivastava gratefully acknowledges the financial assistance from N.B.H.M.(D.A.E) as a Post Doctoral fellowship. The author B.S. Bhadauria is grateful to Banaras Hindu University, Varanasi, India for sanctioning the leave to work



**Fig. 17. Variation of isoconcentrations with time**

as Professor of Mathematics at Department of Applied Mathematics, B. B. Ambedkar University, Lucknow, India.

## REFERENCES

- Alex, S. M., P. R. Patil, and K. S. Venakrishnan (2001). Variable gravity effects on thermal instability in a porous medium with internal heat source and inclined temperature gradient. *Fluid Dynamics Research* 29, 1–6.
- Bejan, A. (1978). Natural convection in an infinite porous medium with a concentrated heat source. *Journal of Fluid Mechanics* 89, 97–107.
- Bhadauria, B. S. (2012). Double diffusive convection in a saturated anisotropic porous layer with internal heat source. *Transport in Porous Media* 92, 299–320.
- Bhadauria, B. S., I. Hashim, and P. G. Siddheshwar (2013). Study of heat transport in a porous medium under g-jitter and internal heating effects. *Transport in Porous Media* 96(1), 21–37.
- Capone, F., M. Gentile, and A. A. Hill (2011). Double diffusive penetrative convection simulated via internal heating in an anisotropic porous layer with throughflow in an anisotropic porous layer with through-flow. *International Journal of Heat and Mass Transfer* 54, 1622–1626.
- Cookey, C. I., V. B. O. Pepple, B. I. Obi, and L. C. Eze (2010). Onset of thermal instability in a low prandtl number fluid with internal heat source in a porous medium. *American Journal of Scientific and Industrial Research* 1, 18–24.
- Govender, S. (2005). Weak non-linear analysis of convection in a gravity modulated porous layer. *Transport in Porous Media* 60(1), 33–42.
- Holzbecher, E. (1998). The influence of variable viscosity on thermal convection in porous media. *Transactions on Engineering Sciences* 20, 115–124.
- Ingham, D. B. and I. Pop (2005). *Transport Phenomena in Porous Media*. Vol. III. Oxford: Elsevier.
- Kuznetsov, A. V. (2006a). Investigation of the onset of bioconvection in a suspension of oxytactic microorganisms subjected to high frequency vertical vibration. *Theoretical and Computational Fluid Dynamics* 20(2), 73–87.
- Kuznetsov, A. V. (2006b). Linear stability analysis of the effect of vertical vibration on bioconvection in a horizontal porous layer of finite depth. *Journal of Porous Media* 9(6), 597–608.
- Kuznetsov, A. V. (2011). The onset of bioconvection in a suspension of negatively geo-tactic microorganisms with high-frequency vertical vibration. *International Communications in Heat and Mass Transfer* 32(9), 1119–1127.

- Kuznetsov, A. V. and D. A. Nield (2008). The effects of combined horizontal and vertical heterogeneity on the onset of convection in a porous medium: double diffusive case. *Transport in Porous Media* 72(2), 157–170.
- Kuznetsov, A. V. and D. A. Nield (2010). The onset of double-diffusive nanofluid convection in a layer of a saturated porous medium. *Transport in Porous Media* 85(2), 941–951.
- Kuznetsov, A. V. and D. A. Nield (2011). Double-diffusive natural convective boundary-layer flow of a nanofluid past a vertical plate. *International Journal of Thermal Sciences* 50(5), 712–717.
- Malashetty, M. S. and V. Padmavathi (1997). Effect of gravity modulation on the onset of convection in a fluid and porous layer. *International Journal of Engineering Science* 35(9), 829–840.
- Nield, D. A. (1996). The effect of temperature-dependent viscosity on the onset of convection in a saturated porous medium. *ASME Journal of Heat Transfer* 118, 803–805.
- Nield, D. A. and A. Bejan (2013). *Convection in Porous Media*, 4th edn. New York: Springer-Verlag.
- Nouri-Borujerdi, A., A. Noghrehabadi, and D. A. S. Rees (2007). Onset of convection in a horizontal porous channel with uniform heat generation using a thermal nonequilibrium model. *Transport in Porous Media* 69, 343–357.
- Nouri-Borujerdi, A., A. Noghrehabadi, and D. A. S. Rees (2008). Influence of darcy number on the onset of convection in porous layer with a uniform heat source. *Inter-national Journal of Thermal Sciences* 47, 1020–1025.
- Payne, L. E., J. C. Song, and B. Straughan (1999). Continuous dependence and convergence for brinkman and forchheimer models with variable viscosity. *Proceedings of the Royal Society of London* 452, 2173–2190.
- Payne, L. E. and B. Straughan (2000). Unconditional nonlinear stability in temperature-dependent viscosity flow in a porous medium. *Studies In Applied Mathematics* 105, 59–81.
- Qin, Y. and J. Chadam (1996). Non-linear convective stability in a porous medium with temperature-dependent viscosity and inertial drag. *Studies in Applied Mathematics* 96, 273–288.
- Razi, Y. P., I. Mojtabi, and M. C. Charrier-Mojtabi (2009). A summary of new predictive high frequency thermo-vibrational modes in porous media. *Transport in Porous Media* 77, 207–208.
- Rees, D. A. S., M. A. Hossain, and S. Kabir (2002). Natural convection of fluid with variable viscosity from a heated vertical wavy surface. *ZAMP* 53, 48–57.
- Rees, D. A. S. and I. Pop (2000). The effect of g-jitter on vertical free convection boundary-layer flow in porous media. *International Communications in Heat and Mass Transfer* 27(3), 415–424.
- Rees, D. A. S. and I. Pop (2001). The effect of g-jitter on free convection near a stagnation point in a porous medium. *International Journal of Heat and Mass Transfer* 44, 877–883.
- Rees, D. A. S. and I. Pop (2003). The effect of large-amplitude g-jitter vertical free convection boundary-layer flow in porous media. *International Journal of Heat and Mass Transfer* 46, 1097–1102.
- Richardson, L. and B. Straughan (1993). Convection with temperature dependent viscosity in a porous medium: non-linear stability and the brinkman effect. *Atti della Accademia Nazionale dei Lincei. Classe di Scienze Fisiche, Matematiche e Naturali. Rendiconti Lincei. Matematica e Applicazioni* 4, 223–232.
- Saravanan, S. (2009). Thermal non-equilibrium porous convection with heat generation and density maximum. *Transp Porous Med* 76, 35–43.
- Siddheshwar, P. G., B. S. Bhadauria, and A. Srivastava (2012a). An analytical study of nonlinear double diffusive convection in a porous medium under temperature/gravity modulation. *Transport in Porous Media* 91, 585–604.
- Siddheshwar, P. G. and A. T. Chan (2004). *Thermorheological effect on Bénard and marangoni convections in anisotropic porous media*. In: Cheng, L., Yeow, K. (eds.) *Hydrodynamics VI Theory and Applications*, Taylor and Francis, London, 471–476.
- Siddheshwar, P. G., R. K. Vanishree, and A. C. Melson (2012b). Study of heat transport in Bénard-Darcy convection with g-jitter and thermo-mechanical anisotropy in variable viscosity liquids. *Transp. Porous Media* 92, 277–288.
- Srivastava, A., B. S. Bhadauria, P. G. Siddheshwar, and I. Hashim (2013). Heat transport in an anisotropic porous medium saturated with variable viscosity liquid under g-jitter and internal heating effects. *Transp Porous Med* 99(1), 359–376.
- Strong, N. (2008). Effect of vertical modulation on the onset of filtration convection. *Journal of Mathematical Fluid Mechanics* 10(4), 488–502.
- Strong, N. (2009). Double-diffusive convection in a porous layer in the presence of vibration. *SIAM Journal on Applied Mathematics* 69(5), 1263–1276.
- Swamy, M. (2014). Effect of g-jitter on the onset of double-diffusive convection in fluid/porous

- layer. *Journal of Porous Media* 17(2), 117–128.
- Swamy, M. S., I. S. Shivakumara, and W. Sidram (2013). The onset of convection in a gravity-modulated viscoelastic fluid-saturated anisotropic porous layer. *Special Topics and Reviews in Porous Media An International Journal* 47(1), 69–80.
- Tveitereid, M. (1977). Thermal convection in a horizontal porous layer with internal heat sources. *International Journal of Heat and Mass Transfer* 20, 1045–1050.
- Vadász, P. (1998). Coriolis effect on gravity-driven convection in a rotating porous layer heated from below. *Journal of Fluid Mechanics* 376, 351–375.
- Vadász, P. (2008). *Emerging Topics in Heat and Mass Transfer in Porous Media*. New York: Springer.
- Vafai, K. (2005). *Handbook of Porous Media*. Boca Raton: Taylor and Francis (CRC).
- Vanishree, R. K. and P. G. Siddheshwar (2010). Effect of rotation on thermal convection in an anisotropic porous medium with temperature-dependent viscosity. *Transport in Porous Media* 81, 73–87.
- Worster, M. G. (1991). Natural convection in a mushy layer. *Journal of Fluid Mechanics* 224, 335–359.